

## Modelling the weather and climate

### Coursework 2: Numerical methods and sub-grid processes

Please hand in your work to the Education Office by 12 noon, Thursday 28 March, 2018. Note: if you hand this assignment in late a mark of 0 will be recorded, unless you make a case for leniency that is accepted by the Mitigation Committee.

Please hand in your workings and MATLAB figures for the questions below. Also hand in fully-commented versions of the MATLAB codes written.

#### 1. Numerical methods

Consider the linear advection equation

$$\frac{\partial \phi}{\partial t} + u \frac{\partial \phi}{\partial x} = 0,$$

on a periodic domain  $0 \leq x \leq 1$ . Consider also the two different initial conditions:

$$\phi(x, 0) = \begin{cases} \frac{1}{2} \{1 + \cos[4\pi(x - 1/2)]\} & \text{if } 1/4 \leq x \leq 3/4 \\ 0 & \text{otherwise} \end{cases} \quad (1)$$

$$\phi(x, 0) = \begin{cases} 1 & \text{if } 1/4 \leq x \leq 3/4 \\ 0 & \text{otherwise} \end{cases} \quad (2)$$

- (a) Write a MATLAB code to integrate the linear advection equation until  $t = 0.5$  using a centred-in-time, centred-in-space (CTCS) discretization. Perform the integration for both sets of initial conditions. Take  $u = 2.0$ , use  $N = 51$  equally spaced gridpoints, and take  $\Delta t = 0.005$ . You might find it helpful to follow the structure of `diffuse.m` (available on the ele module web-page) and use the same tricks for coping with the periodic boundary condition. (30)
- (b) Discretise the advection equation using a semi-Lagrangian scheme, and cubic-Lagrange interpolation. The departure points can be evaluated using a simple linear formula, i.e. a displacement of  $u\Delta t$  upstream. Demonstrate how this formulation improves on the CTCS scheme when increasing the time step above the CFL limit. (35)

#### 2. Sub-grid processes

Holton, section 5.3 gives a mixed-layer model of the atmospheric boundary layer

$$\bar{u} = \bar{u}_g - \kappa_s |\bar{\mathbf{V}}| \bar{v}, \quad \bar{v} = \kappa_s |\bar{\mathbf{V}}| \bar{u},$$

where the turbulence magnitude is given by  $\kappa_s = C_d/(fh)$ .

- (a) Explain the relationship between the above mixed-layer equations and the Reynolds averaged momentum equations, explaining carefully the meaning of the notation. (3)

(b) You are given the code `mixed_layer_wind1.m`. Modify the code to calculate the wind turning (the angle between the geostrophic and mixed layer winds) for a range of values of  $h$ . Use the values:  $u_g = 10 \text{ ms}^{-1}$ ,  $C_d = 2 \times 10^{-3}$  and  $f = 10^{-4} \text{ s}^{-1}$ . Calculate and plot the wind turning for values of  $h$  in the range of 300 m to 3000 m.

(12)

(c) You are given the code `mixed_layer_wind2.m` which prescribes the geostrophic wind ( $u_g, v_g$ ) at each horizontal point. Calculate the mixed-layer wind ( $\bar{u}, \bar{v}$ ) at each horizontal point using the algorithm in `mixed_layer_wind1.m`. *Hint: you may want to incorporate the mixed-layer algorithm as a function.* Plot the mixed-layer wind as a vector-wind plot (as done for the geostrophic wind). Assuming a constant boundary-layer depth of 1000 m, calculate the Ekman pumping for each horizontal point and plot the results.

(20)

[100]